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# 1 **Mangrove mortality in a changing climate: An** 2 **overview**

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19 **Abstract**

20 *Mangroves provide vital ecosystem services at the dynamic interface between land and*  
21 *oceans. Recent reports of mangrove mortality suggest that mangroves may be adversely*  
22 *affected by climate change. Here, we review historical mortality events from natural causes*  
23 *(all mortality other than deforestation, land use change and pollution) and provide a global*  
24 *assessment of mortality drivers. Since the 1960's approximately 36,000 ha of mangrove*  
25 *mortality has been reported (0.2% of total mangrove cover in 2011) in 47 peer reviewed*  
26 *articles. Due to the uneven global distribution of research effort, it is likely that mangrove*  
27 *mortality events go unreported in many countries. It is therefore difficult to assess temporal*  
28 *changes in mortality due to the small number of reports and increasing effort in observations*  
29 *in recent years. From the published literature, approximately 70% of reported mangrove loss*  
30 *from natural causes has occurred as a result of low frequency, high intensity weather events,*  
31 *such as tropical cyclones and climatic extremes. Globally, tropical cyclones have caused the*  
32 *greatest area of mangrove mortality, equivalent to 45% of the reported global mangrove*  
33 *mortality area from events over six decades. However, recent large-scale mortality events*  
34 *associated with climatic extremes in Australia account for 22% of all reported historical*  
35 *forest loss. These recent mortality events suggest the increasing importance of extreme*  
36 *climatic events, and highlight that mangroves may be important sentinels of global climate*  
37 *change. Increasing frequency, intensity and destructiveness of cyclones as well as climatic*  
38 *extremes, including low and high sea level events and heat waves, have the potential to*  
39 *directly influence mangrove mortality and recovery, particularly in mid latitudes.*

40

## 41 1.1 Introduction

42 Mangrove forests provide a wide range of highly valuable ecosystem services, including  
43 support of biodiversity and fisheries, coastal protection, carbon sequestration and nutrient  
44 processing (Barbier et al., 2011). These forests sequester large amounts of carbon from the  
45 atmosphere into their biomass and soils (Mcleod et al., 2011), and contribute to marine food  
46 webs via the detrital energy flow pathways (Lugo et al., 1974, Dittmar et al., 2006, Abrantes  
47 et al., 2014). The collective economic valuation of mangrove ecosystem services has been  
48 estimated to be worth USD 194,000 per hectare per year, with a global value of USD 2.7  
49 trillion per year (Barbier et al., 2011, Costanza et al., 2014).

50 Despite their importance, extensive deforestation of mangroves has occurred (Hamilton et al.,  
51 2016, Richards et al., 2016). Further, anthropogenic activities such as changes in hydrology,  
52 add synergistic stressors to mangrove forests, reducing ecosystem function and ecosystem  
53 service value (Lewis et al., 2016). Although deforestation has been the major cause of forest  
54 loss in the past, mangrove loss from various natural and anthropogenic disturbances have also  
55 been reported in the literature (Jimenez et al., 1985). Mangroves are also susceptible to  
56 climate change, which has been anticipated to have widespread negative consequences for  
57 mangrove distribution (Lovelock et al., 2007, Alongi, 2008, Gilman et al., 2008). However,  
58 few studies have unequivocally demonstrated the effects of climate change on mangroves.

59 Recent reports of mangrove mortality in Australia (Duke et al., 2017, Lovelock et al., 2017)  
60 and elsewhere (Albert et al., 2016, Servino et al., 2018) suggest that the incidences of  
61 mangrove mortality may be increasing due to climate change. However, most research on the  
62 effects of climate change on mangroves has focused on temperature driven effects such as  
63 poleward expansion and sea level rise (Gilman et al., 2008, Lovelock et al., 2015, Kelleway  
64 et al., 2017, Osland et al., 2017a). A global analysis of mangrove mortality has not been  
65 performed since Jimenez et al. (1985) collated reports of mangrove mortality three decades  
66 ago. Here we review global records of mangrove losses from natural events, and analyse the  
67 environmental factors causing these losses. We consider the impacts of climate change to  
68 mangroves to be a natural phenomenon since we cannot distinguish individual climatic events  
69 as natural or anthropogenic. However, it is clear that anthropogenic activities which induce  
70 change to the global climate will influence individual climatic events and the future  
71 mangrove distribution (Harris et al., 2018, Lovelock and Ellison 2007, Alongi 2008, Gilman  
72 et al., 2008).

73

## 74 2.1 Historical mortality reports and emerging trends

75 Here we differentiate between mangrove forests which have undergone disturbance yet have  
76 the ability to recover (dieback) and the death of a forest related to a disturbance (mortality)  
77 which does not have the ability to recover in the existing environmental conditions (Table 1).  
78 We also differentiate between tree mortality (individual level) and forest mortality  
79 (ecosystem level). Forest and individual tree recovery can include processes such as re-  
80 sprouting (coppicing) or regeneration via propagule recruitment. The time frame of this  
81 recovery may range from years to decades, depending upon the nature of the mortality event  
82 and the resilience of the forest. Inability to recover can be caused by permanent changes in  
83 environmental conditions that prevent recovery (e.g. subsidence, erosion or permanent  
84 changes in hydrology). Mangrove mortality from natural causes is reported for every  
85 continent where mangroves are found, however the highest frequency of mangrove mortality  
86 is reported from the Caribbean and the Gulf of Mexico (Figure 1; Table 2). To our

87 knowledge, excluding deforestation, forest mortality areas exceeding 5,000 ha have only been  
 88 reported from Australia Duke et al. (2017). Considering the uneven global distribution of  
 89 research effort, it is likely that mangrove mortality events may go unreported in many  
 90 countries.

91 **Table 1. Definitions of terms**

<b>Dieback</b>	Canopy loss that may or may not lead to forest mortality.
<b>Mortality</b>	Forest death caused by disturbance and/or change in environmental conditions.
<b>Climatic extremes</b>	Climatic events such as drought, flood, extreme heat or cold, extreme sea level variation, localized weather events and intense wind or wave energy. Defined by Smith (2011) as the “occurrence in which a statistically rare or unusual climatic period which has potential to alter ecosystem structure and/or function well outside the bounds of what is considered typical or normal variability”. The response of mangroves to these climatic extremes is site and species specific.
<b>Extreme climatic event</b>	An extreme ecosystem response to climatic extremes, beyond the ecosystems normal range of variability (Smith, 2011).
<b>Sea level variability</b>	Short term (days to months) increase or decrease in regional sea level caused by atmospheric pressure, winds, air/sea temperature, freshwater inputs and ocean currents.
<b>Sea level rise</b>	Long-term global increase in sea level.

92



93  
 94 **Figure 1. Global mangrove area along with the location and extent of mangrove mortality events driven**  
 95 **by natural causes. Image of global mangrove area is modified from (Giri et al., 2011).**

96 Over short time scales, mangrove forests experience daily tide fluctuations, which can  
 97 significantly influence water availability, soil salinity and anoxia, with a direct effect upon  
 98 plant growth and fitness (Ball, 1988). Over longer time scales, mangroves have experienced  
 99 large displacements in their distribution due to historical sea level change over the Holocene  
 100 (Woodroffe et al., 1991) and over shorter time scales (years, decades) due to varying  
 101 sediment supply (Woodroffe et al., 2016). Mangroves can undergo rapid contraction in cover  
 102 when conditions are suboptimal (Eslami-Andergoli et al., 2015) or expansion under  
 103 favourable conditions (Asbridge et al., 2016). On a global scale, mangrove mortality due to  
 104 natural causes has been reported over scales ranging from 1 – 7,400 hectares (Figure 1; Table

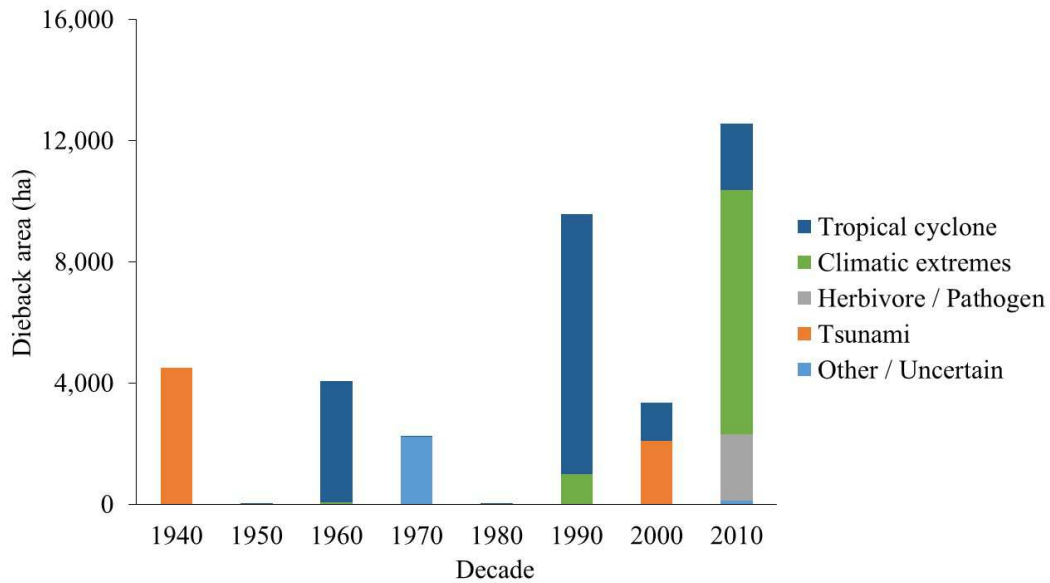
105 2) with a total area of forest mortality reported since the 1960s of approximately 36,000 ha  
 106 (0.2% of total mangrove cover in 2011).

107 **Table 2. Summary of reported mangrove mortality from natural causes reported in peer reviewed**  
 108 **literature expanded from Jimenez et al. (1985). For review of mortality events in plantations and restored**  
 109 **forests see López-Portillo et al. (2017).**

Site	Date	Cause	Area (ha)	Source
<b>Tropical cyclones</b>				
Florida, USA	1935	Labor Day Hurricane	Extensive	(Wanless and Vlaswinkel, 2005, Smith et al., 2009)
Puerto Rico	1956	Hurricane Betsy	< 5*	(Wadsworth et al., 1959)
Florida, USA	1960	Hurricane Donna	~4,000**	(Craighead et al., 1962)
Florida, USA	1960	Hurricane Donna	Patches	(Craighead, 1964)
Belize	1961	Hurricane Hattie	Patches	(Vermeer, 1963)
Australia	1971	Cyclone Althea / siltation	~5*	(Heinsohn et al., 1974)
Nicaragua	1988	Hurricane Joan	~ 5*	(Roth, 1992)
Florida, USA	1992	Hurricane Andrew	150	(Ogden, 1992)
Dominican Republic	1998	Hurricane Georges	2,240	(Sherman et al., 2001)
Honduras	1998	Hurricane Mitch	490*	(Cahoon et al., 2003)
Australia	1999	Cyclone Vance	5,700	(Paling et al., 2008)
Micronesia	2004	Cyclone (Typhoon) Sudal	< 10*	(Kauffman et al., 2010)
Florida, USA	2005	Hurricane Wilma	1,250	(Smith et al., 2009)
Philippines	2009	Typhoon Chan- hom	Extensive	(Salmo et al., 2014)
Australia	2011	Cyclone Yasi	2200	Asbridge et al., (in review)
<b>Climatic extremes</b>				
Puerto Rico	1937- 1972	Hypersalinity	Fringes	(Cintrón et al., 1978)
South Africa	1965	Flooding	~ 10	(Breen et al., 1969)
Florida, USA	1983	Freeze	~1	(Montague et al., 1997)
Gulf of Mexico, USA	1930 - 1980	Freeze	Extensive	(Sherrod and McMillan, 1981)
Gulf of Mexico, USA	1983	Freeze	Extensive	(Sherrod and McMillan, 1985)
Guadeloupe	-	Hypersalinity	Extensive	(Servant et al., 1978) as cited in Jimenez et al. (1985)
Java	-	Flooding/siltation	42	(Soerianegara, 1968) as cited in Jimenez et al. (1985)
Florida, USA	-	Frost	Extensive	(Lugo et al., 1977) as cited in Jimenez et al. (1985)
Texas	-	Frost	Patches	(West, 1977) as cited in Jimenez et al. (1985)
Gulf of Mexico, USA	1983 and	Freeze	Extensive	(Lonard and Judd, 1991)

	1989			
Florida, USA	1980's	Freeze	15	(Stevens et al., 2006)
Senegal	1990's	Drought	Extensive	(Diop et al., 1997)
Australia	1994	Hail storm	0.1	(Houston, 1999)
Venezuela	1990 - 1998	ENSO related drought	880	(Barreto, 2008, Otero et al., 2017)
Tanzania	1997	Flooding	117	(Erfemeijer et al., 2005)
Kakadu, Australia	2015 - 2016	ENSO related climatic extremes	Extensive	(Lucas et al., 2018)
Mangrove Bay, Australia	2002-2003, 2015 - 2016	ENSO related climatic extremes	40	(Lovelock et al., 2017)
Gulf of Carpentaria, Australia	2015-2016	ENSO related climatic extremes	~7,400	(Duke et al., 2017)
Brazil	2016	Hail storm	~500	(Servino et al., 2018)
<b>Tsunamis</b>				
Dominican Republic	1946	Dominican Republic Tsunami	~ 4,500	(Sachtler, 1973) as cited in Jimenez et al. (1985)
Indonesia	2004	Banda Aceh, Tsunami	300 – 750	(BAPPENAS, 2005)
Thailand	2004	Banda Aceh, Tsunami	1,050	(Kamthonkiat et al., 2011)
Great Nicobar Island, India	2004	Banda Aceh, Tsunami	530	(Sridhar, 2007)
Thailand	2004	Banda Aceh, Tsunami	patches	(Fujioka et al., 2008)
Thailand	2004	Banda Aceh, Tsunami	~1*	(Yanagisawa et al., 2010)
Andaman Islands	2004	Banda Aceh, Tsunami	Patches	(Roy, 2016)
<b>Herbivores and pathogens</b>				
Australia	1980	Pathogen	Patches	(Weste et al., 1982)
Malaysia	1990's	Herbivory	Patches	(Jin-Eong, 1995)
India	2000	Herbivory	Patches	(Kathiresan, 2003)
Kenya	2005	Herbivory	Patches	(Jenoh et al., 2016)
Mexico	2010	Herbivory	3,846	(Sánchez et al., 2018)
<b>Other</b>				
Bahamas	2011	Cumulative stressors	120	(Rossi, 2018)
The Gambia	1970's	Uncertain	2,211	(Blasco, 1983)
Bermuda	1800's - 1993	Sea level rise	1	(Ellison, 1993)

111 Assessment of mangrove mortality over time (decadal increments) indicates a general  
 112 increase in mortality area each decade since the 1940's (Figure 2). Climatic extremes and  
 113 tropical cyclones account for ~80 % of the mangrove mortality area reported since 1990. The  
 114 increasing mangrove mortality area observed each decade may be due to increasing research  
 115 effort and advances in remote sensing technology, thus the trend of increasing mortality area  
 116 observed here may not necessarily be related to increasing mangrove mortality, but may  
 117 represent a research effort bias. When the area of mortality is adjusted for research effort  
 118 (standardized using the number of publications returned using a search on "mangrove" in the  
 119 Web of Science for each decade), no significant trend is observed (not shown).



120

121 **Figure 2. Summary of decadal trends of total global area of mangrove mortality from various natural**  
 122 **causes reported in peer reviewed literature summarized in Table 2.**

123 The global distribution, abundance and species richness of mangrove forests is controlled by  
 124 climatic parameters, predominantly rainfall and temperature (Osland et al., 2017b). However,  
 125 these climatic parameters are strongly influenced by regional scale processes. Therefore,  
 126 forest distributions are often largely controlled by regional and local scale environmental  
 127 controls such as temperature, precipitation, tidal amplitude, wave energy and riverine inputs  
 128 (Bunt, 1996, Alongi, 2009). Natural events that cause mangrove mortality can result from  
 129 individual and/or synergistic changes in environmental conditions that exceed mangrove  
 130 physiological limits (Table 3).

131 **Table 3. Causes of mangrove mortality or dieback from natural events.**

Disturbance	Cause of forest mortality or dieback	References
<b>Tropical cyclones</b>	Strong winds, flooding rain, high-energy waves and storm surges. Effects include physical damage to canopies and soils (erosion); and sedimentation and flooding which limit oxygen supply to roots.	(Paling et al., 2008, Smith et al., 2009, Kauffman et al., 2010)
<b>Heat waves/frost</b>	At temperatures exceeding 38-40°C, photorespiration increases while photosynthesis is inhibited. Evaporation increases with high temperatures which increases water loss and risk of desiccation. Frost /freezes result in tissue damage and hydraulic failure at high latitudes.	(West 1977) (Clough et al., 1982, Rennenberg et al., 2006, Stuart et al., 2007,



		Lindner et al., 2010)
<b>Drought</b>	Drought stress is directly related to the development of hypersaline soil conditions. Low rainfall and groundwater inputs reduce water availability and high temperatures increase evaporation and evapotranspiration resulting in hydraulic failure and desiccation.	(Medina and Francisco 1997, Hoppe-Speer et al., 2013, Ward et al., 2016)
<b>Tsunamis</b>	Extreme physical damage to canopies and soils (erosion). Sedimentation and flooding which limit oxygen supply to roots.	(Kamthonkiat et al., 2011; Sachtler, 1973
<b>Herbivores and pathogens</b>	Pathogens - leaf damage and wood decomposition in living plants, including symptoms such as butt, heart and root rot. Herbivores – tree defoliation and loss of photosynthetic function.	(Hyde et al., 1998). (Osorio et al., 2016)

132

133 **2.1.1 Tropical cyclones** – Extreme storms have caused the greatest global area of mangrove  
134 mortality, equivalent to 45% of the reported global mangrove mortality area (Table 2; Figure  
135 2). In the last five decades the percentage of mangrove mortality caused by tropical cyclones  
136 has remained relatively constant, accounting for 30-60% of the total mortality area (Figure 3).  
137 Tropical cyclones generate strong winds, flooding rain, high-energy waves and storm surges  
138 that can negatively affect coastal wetlands (Table 3). Individual events have caused damage  
139 to areas up to 28,000 ha (Villamayor et al., 2016) and forest mortality in areas of up to 5,700  
140 ha (Table 2).

141 Storm trajectories have high inter-annual variability, yet broad cyclone paths exist due to  
142 oceanic boundary conditions and sea surface temperatures (Ulbrich et al., 2009). For  
143 example, the Gulf of Mexico is frequently affected by hurricanes which have caused  
144 considerable mangrove damage and mortality (Table 2, Craighead, 1964, Smith et al., 2009).  
145 Indeed, coastal ecosystems along the Gulf of Mexico have been subject to one major  
146 hurricane every 3 years over the last century (Diaz et al., 2012). Cumulative impacts from  
147 reoccurring mortality events related to hurricane frequency are not well understood, but can  
148 clearly influence recovery and may be related to permanent ecosystem changes in some cases  
149 (Smith et al., 2009, Feller et al., 2015).

150 The northern hemisphere experiences ~70% of global cyclones/hurricanes (Knutson et al.,  
151 2010) and 76% of reported cyclone mortality events (Table 2), yet these reports only account  
152 for 51% of the global cyclone related mortality area (Table 2). Nearly 100% of cyclone  
153 associated mangrove mortality reported in the Southern hemisphere has occurred in Australia,  
154 perhaps due to larger research effort. The relative importance of cyclone associated mortality  
155 has declined in the last decade as mortality associated with climatic extremes has increased  
156 (Figure 3).

157 **2.1.2 Climatic extremes** – In previous centuries, mangrove coverage at high latitudes has  
158 been controlled by extreme cold events causing frost/freezes (Sherrod et al., 1981, Sherrod et  
159 al., 1985, Osland et al., 2017a). This is still the case in many continents, however poleward  
160 expansion of mangroves, and encroachment into saltmarsh is already occurring as global

161 temperatures increase (Saintilan et al., 2014). Prior to the 1980's, mangrove forest mortality  
162 had not been attributed to climatic extremes other than frost/freeze events (Table 2). In the  
163 past decade, extreme climatic events have accounted for 22% of the reported global forest  
164 loss (Figure 3) and 98% of this mortality have been due to drought and regional sea level  
165 variability. El Niño Southern Oscillation (ENSO) are large scale weather patterns influencing  
166 seasonal rainfall and sea levels (Moon et al., 2015, Widlansky et al., 2015). ENSO driven low  
167 sea level and high temperatures have severe impacts on nearshore ecosystems and the  
168 reoccurrence intervals of these events may be critical for slow recovering coastal ecosystems  
169 such as mangroves (Glynn, 1988). Strong ENSO events will clearly play an important role in  
170 mangrove mortality events into the future.

171 In Australia and the Eastern Pacific, mangrove mortality associated with drought is  
172 overwhelmingly associated with ENSO events. A recent mortality event in the Gulf of  
173 Carpentaria (northern Australia) is unprecedented in areal extent and was attributed to the  
174 combined climatic extremes associated with a severe El Niño event (Duke et al., 2017, Harris  
175 et al., 2017). Other recent mangrove forest mortalities due to extremes in temperature,  
176 drought, localised weather events and regional sea level variability have occurred in Exmouth  
177 in Western Australia, Kakadu in North Western Australia, the Sundarbans and Brazil (Paul et  
178 al., 2017, Lovelock et al., 2017, Lucas et al., 2018, Servino et al., 2018).

179 Prior to these recent events, few studies had observed extensive drought related mortality  
180 (Figure 2). El Niño associated drought caused 880 ha of mangrove forest mortality in  
181 Venezuela during a period from 1990 – 1994 and during an intense El Niño period from 1997  
182 – 1998 (Table 2, Barreto, 2008, Otero et al., 2017). In these cases, mangrove forests showed  
183 no recovery nine years later, due to peat oxidation and sediment subsidence that  
184 fundamentally changed the suitability of the site for mangrove growth (Otero et al., 2017). El  
185 Niño associated drought also caused substantial stress to mangrove forests in Micronesia,  
186 with reports of a two fold increase in porewater salinity, although no tree mortality was  
187 reported (Drexler et al., 2001). Severe drought in the Eastern Pacific associated with La Niña  
188 has also caused a massive mortality of salt marsh in 2000, affecting an area of 100,000 ha  
189 equivalent to 41% of contiguous coastal wetlands in the United States (McKee et al., 2004).

190 **2.1.3 Tsunamis** - The frequency of tsunamis is approximately 10% that of extreme storm  
191 events (Intergovernmental Oceanographic Commission 2016; Webster 2005). However,  
192 tsunamis have accounted for 18% of historical mangrove mortality and have historically  
193 accounted for some of the largest recorded mangrove mortality events (Figure 2). The 2004  
194 Indian Ocean earthquake triggered a series of devastating tsunamis that resulted in the loss of  
195 1880 – 2,330 ha of mangroves, as well as massive damage to coral reefs, sand dunes, forests,  
196 groundwater reserves, and to human settlements and agricultural lands (Kamthonkiat et al.,  
197 2011). GIS analysis of forests damaged by the 2004 tsunami revealed that approximately half  
198 of affected forest areas experienced total destruction, while half experienced severe damage  
199 that removed most trees (Kamthonkiat et al., 2011).

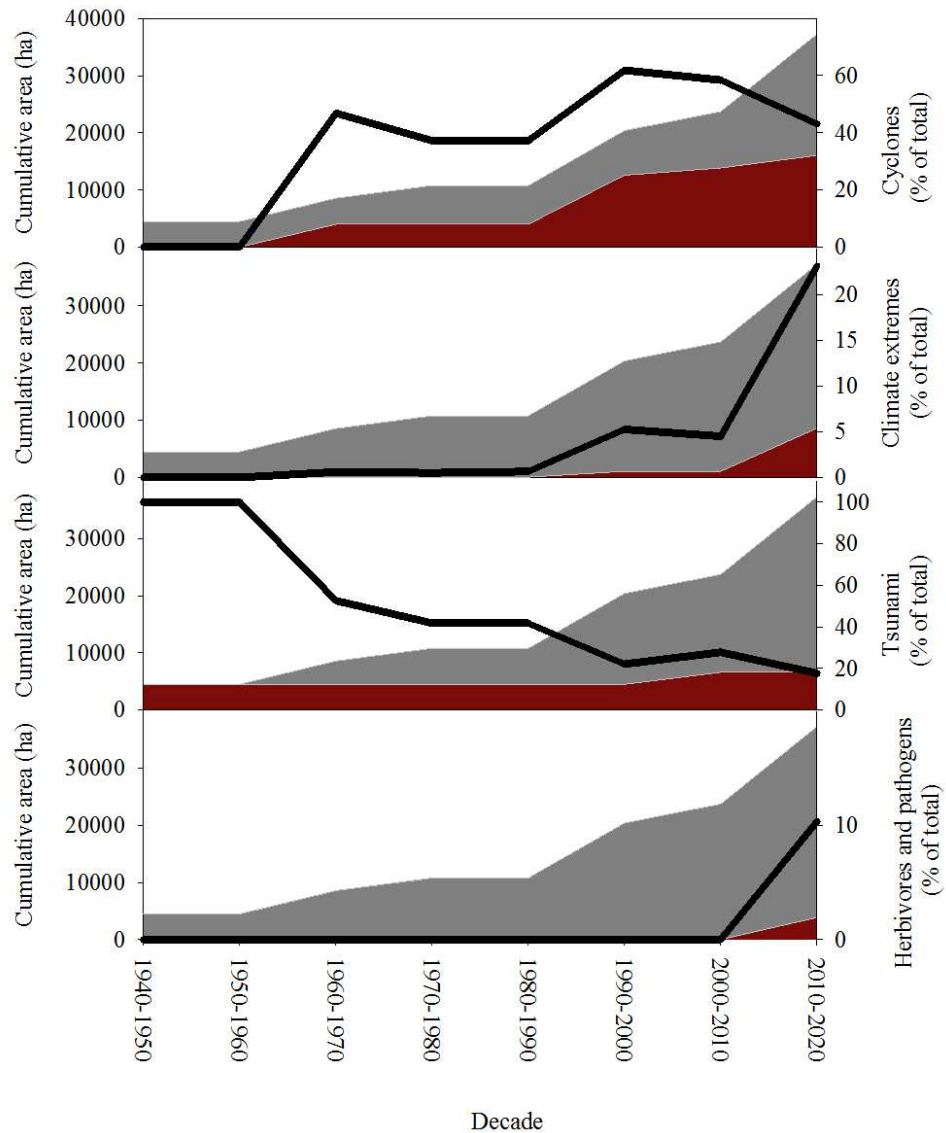
200 The first report of a major mangrove mortality caused by a tsunami was in 1946 in the  
201 Dominican Republic, which caused a mortality area estimated at 4,500 ha (Sachtler, 1973,  
202 Jiminez et al., 1985). Since this event, only one other significant forest loss associated with  
203 tsunami has been reported (Indian Ocean 2004 tsunami). Because tsunami related dieback is  
204 uncommon and sporadic, the apparent decrease in total mangrove mortality area over time

205 caused by tsunamis should be interpreted with care, since other natural causes have caused  
206 increasingly extensive areas of forest mortality (Figure 3).

207 **2.1.4 Herbivores and pathogens** – Historically, while several works document insect  
208 infestations in mangroves (e.g. Feller and Mathis 1997, Feller and McKee, 1999), few cases  
209 have reported significant forest loss from herbivores and pathogens. Cases of tree mortalities  
210 caused by wood borer insects (Feller, 2002), defoliation (Anderson et al., 1995, Jin-Eong,  
211 1995, Mehlig et al., 2005) and pathogens (Pegg et al., 1980) have been reported, but only in  
212 small, isolated patches. Jenoh et al. (2016) describe serious damage to extensive areas of  
213 mangrove forests in Kenya from two species of wood boring insects, which caused patches of  
214 forest dieback. These smaller scale losses of forest area or canopy may cause a reduction in  
215 ecosystem services by changing forest structure and function (Jennerjahn et al. 2017), which  
216 can also decrease forest resilience.

217  
218 Recently, Sánchez et al. (2018) report the first case of extensive forest mortality from  
219 herbivory in Mexico 2010 (Table 2: Figure 3). In this case a mangrove forest area of 3846 ha  
220 was severely damaged from defoliation caused by an overpopulation of the caterpillar  
221 *Anacamptodes sp.*, resulting in an area of 2,196 ha of forest mortality (Sánchez et al., 2015,  
222 Sánchez et al., 2018). Herbivores and pathogens may stress forests and cause forest mortality  
223 in combination with other stressors (Rossi, 2018).

224



225

226 **Figure 3. Total cumulative mortality area (grey), cumulative areas (red) of a) cyclones, b) climate**  
 227 **extremes, c) tsunamis and d) herbivores and pathogens. The percentage of mortality from each individual**  
 228 **cause of the total mortality area per decade (black line). Mortality area from 49 peer reviewed reports.**

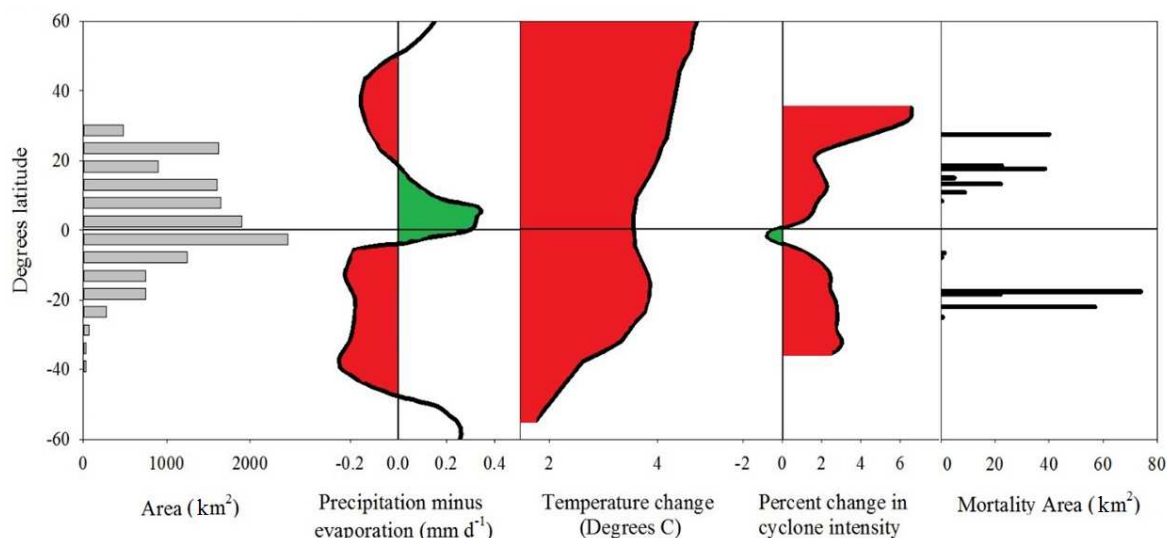
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### 230 3.1 Future outlook

231

232 **3.1.1 Tropical cyclones**– Cyclone intensity, destructiveness and frequency has increased  
 233 since the 1970's (Hartmann et al., 2013). While this trend may have large regional  
 234 differences (Knutson et al., 2010, Wong et al., 2014), cyclone intensity is predicted to  
 235 continue to increase in areas where mangrove mortality is already occurring (Figure 4). These  
 236 climatic projections compared with historic mortality events suggest an increasing threat to  
 237 non-equatorial mangroves, which may be increased in areas affected by human activities, for  
 238 example nutrient enrichment (Feller et al. 2015). Intensified temporal clustering of cyclones  
 239 is considered a major contributor to coral reef decline (Wolff et al., 2016) and could have  
 240 similar impacts on mangrove ecosystems through reduction in recovery time, although they  
 241 may also enhance dispersal of mangrove propagules (Dangremond et al., 2016) and delivery

242 of nutrients and freshwater (Adame et al., 2011). Modelling suggests that cyclones will have  
 243 enhanced precipitation and increased latitudinal range into the future due to climate change,  
 244 but there are high levels of uncertainty in these models (Wong et al., 2014, Scoccimarro et  
 245 al., 2017).



246

247 **Figure 4. Projected or measured influence of global climate change on climatic variables. a) Mangrove**  
 248 **area (Giri et al., 2011), b) mean predicted precipitation minus evaporation ( $\text{mm d}^{-1}$ ) over land and ocean**  
 249 **(Byrne et al., 2015), c) predicted temperature change ( $^{\circ}\text{C}$ ) for 2080–2099 (Meehl et al., 2007), d) modelled**  
 250 **poleward shift in cyclone activity measured between 1980–1994 to 1995–2010 (Kossin et al., 2014) and e)**  
 251 **The area and latitude of individual mangrove mortality events associated with climatic disturbance as**  
 252 **reported in peer reviewed literature (See Table 2).**

253 **3.1.2 Climatic extremes** – The increasing trend in mangrove mortality due to climate  
 254 extremes (Figure 3) may be associated with climate change and could increase as climate  
 255 change progresses. However this may be counterbalanced to some extent by increasing  
 256 poleward expansion (Osland et al., 2017a). For example, extensive losses of mangrove  
 257 coverage have occurred historically due to periods of subfreezing temperature (Sherrod et al.,  
 258 1981, Sherrod et al., 1985). Reduction in frequency of freezing events is already allowing for  
 259 range expansion of mangroves to higher latitudes (Osland et al., 2017a). In addition,  
 260 increasing precipitation in low latitudes may benefit equatorial mangrove forests (Sanders et  
 261 al., 2016). However, increased precipitation may also lead to increased  
 262 siltation/sedimentation, which could have adverse effects (Ellison, 1999). Projected changes  
 263 to precipitation, evaporation and temperature will also result in reduced water availability to  
 264 non-equatorial mangroves (Figure 4).

265 Increased mortality due to aridity driven by climate change was predicted for Central  
 266 America, the Caribbean and North Western Australia (Record et al., 2013, Alongi, 2015).  
 267 The intensity of El Niño events in the Pacific Ocean has increased by approximately 50% in  
 268 the last 30 years (Lee et al., 2010, Cai et al., 2014). This is increasing the frequency and  
 269 intensity of climatic extremes and is leading to intensified sea level variability (Cai et al.,  
 270 2012, Moon et al., 2015). In particular, extreme low sea level events are predicted to double  
 271 in occurrence (Widlansky et al., 2015). Low sea levels are associated with high atmospheric  
 272 pressure, and thus are concurrently associated with low humidity and rainfall, which may  
 273 increase the likelihood of mangrove mortality in mid-latitude ENSO influenced regions.

274 Low sea levels which result in hypersaline soil water in combination with low humidity and  
275 drought is a severe threat to mangroves due to their synergistic adverse effects on  
276 physiological functioning (Lovelock et al., 2017). These environmental conditions are  
277 implicated in the recent mortality events in Northern and Western Australia, which account  
278 for ~20% of globally reported mortality (Duke et al., 2017, Harris et al., 2017, Lovelock et  
279 al., 2017, Lucas et al., 2018). These observed mortality events are consistent with the  
280 prediction that mangroves in the Caribbean, Central America, West and South Africa, the  
281 Persian Gulf, the Red Sea and Australia are particularly susceptible to mortality from  
282 increasing aridity under future climate change scenarios (Alongi, 2015, Jennerjahn et al.,  
283 2017).

284 **3.1.3 Sea level rise, wave energy and sediment availability** - Sea level rise is predicted to  
285 result in mangrove mortality in areas with low sediment accretion (Gilman et al., 2008), or  
286 where there are barriers to mangrove transgression landward (i.e. the coastal squeeze effect,  
287 Torio and Chmura, 2013). Yet there are few reports of sea level rise-driven mortality which  
288 are not associated with land-use change or other direct human influence. Sediment supply is  
289 critical to enable vertically accretion, thereby allowing mangroves to keep pace with sea-level  
290 rise. (Kirwan et al., 2013, Krauss et al., 2014, Lovelock et al., 2015, Woodroffe et al., 2016)  
291 and thus reductions in sediment supply may lead to mangrove mortality. Sediment delivery to  
292 the coast is strongly influenced by river flows. Therefore, damming of rivers and other  
293 anthropogenic activities decreasing sediment supply to coastal ecosystems (Milliman et al.,  
294 2013, Wong et al., 2014), as well as climate change associated reductions in river flows,  
295 could contribute to future mortality as resilience of mangroves is reduced.

296  
297 Mangroves are also highly sensitive to wave energy (Barnard et al., 2015, Walcker et al.,  
298 2015, Albert et al., 2016). Waves can drive mortality due to erosion (Thampanya et al.,  
299 2006). Increasing wave heights associated with climate change have already been attributed  
300 to severe shoreline recession in the Solomon Islands (Albert et al., 2016) and large scale loss  
301 of mangrove forest area due to erosion have been observed with periods of increased wave  
302 energy during positive cycles of the North Atlantic Oscillation winter index (Walcker et al.,  
303 2015). Waves also reduce recruitment through dislodgement of seedlings (Balke et al., 2013),  
304 which can limit recovery after mortality events. Oceanic wind speeds have increased due to  
305 climate change (Wong et al., 2014), which affects wave heights, current regimes and  
306 upwelling systems (Narayan et al., 2010, Albert et al., 2016), all factors that could contribute  
307 to future mangrove mortality.

308  
309 **3.1.4 Multiple stressors** - Several recent examples highlight mangrove loss associated with  
310 cumulative impacts from multiple stressors that cannot be separated. The largest reported  
311 area of mangrove tree mortality from herbivory occurred in the Mexico in 2010 (Sánchez et  
312 al., 2018). The mortality occurred in an area highlighted as at risk from cumulative impacts of  
313 stressors such as erosion and salt water intrusion (Sánchez et al., 2018). The gradual decline  
314 of mangroves to cumulative impacts of herbivores, pathogens and climatic extremes  
315 highlighted in a recent study in the Bahamas (Rossi, 2018), provides an example of how  
316 multiple minor stressors can cumulatively result in mortality to mangrove forests. Similar  
317 large scale dieback associated with combined climatic stressors and herbivory have also been  
318 reported in salt marsh (Silliman et al., 2005, He et al., 2017). In another recent example,  
319 Servino et al. (2018) report a large scale mangrove mortality from a hailstorm occurring in  
320 forests stressed by El Niño associated drought over multiple years. As such, it is important to  
321 consider the synergistic effect of multiple stressors when assessing the vulnerability of  
322 mangrove forests to climate change.

323 Climate change has the potential to influence mangrove mortality through the complex  
324 synergistic changes in environmental parameters. In terrestrial forests, herbivory and disease  
325 may cause forest mortality in conjunction with the many cumulative stressors associated with  
326 climate change (Williams et al., 1995, Lindner et al., 2010, Young et al., 2017). Future  
327 mangrove mortality will be influenced by both direct and indirect interactions among global  
328 sea level rise, regional sea level variability, global change in temperature, evaporation and  
329 precipitation, increased cyclone latitudinal range, intensity and frequency and changes to  
330 wind energy, wave energy and sediment budgets. In addition there will be further complexity  
331 associated with interactions between natural and anthropogenic processes (Figure 4, Alongi,  
332 2015, Feller et al., 2015, Duke et al., 2017, Lovelock et al., 2017).

333

#### 334 **4.1 Conclusions**

335 Half of the historic 36,148 ha reported loss of mangrove forest from natural causes resulted  
336 from tropical cyclones. Increases in the frequency and intensity of storms may therefore be  
337 the greatest natural threat to mangroves in the future, particularly where this interacts in  
338 synergy with high rates of relative sea level rise (Lovelock et al., 2015, Albert et al., 2016,  
339 Ward et al., 2016), which can be exacerbated by reduced sediment supply and subsidence  
340 (Lovelock et al., 2015). Landward migration and poleward expansion, where possible, may  
341 reduce impacts on the extent of mangrove cover (Traill et al., 2011, Rogers et al., 2012,  
342 Osland et al., 2017b).

343 Since 2015, widespread areas of forest mortality have been linked to climatic extremes such  
344 as drought and sea level variability, which are now accounting for a rapidly increasing  
345 proportion of global forest loss. This review highlights the emerging importance of increasing  
346 extreme climatic events and variation in frequency and intensity of El Nino events in driving  
347 mangrove mortality (Duke et al., 2017, Lovelock et al., 2017, Otero et al., 2017). Therefore,  
348 increasing synergistic impacts of intense storms and climate extremes may reduce resilience  
349 of mangroves and the ecosystem services they provide, particularly in subtropical latitudes.

350 It is important to note that the complex and interrelated changes to climatic variables may not  
351 lead to predictable environmental responses. For example, increased precipitation in low  
352 latitudes may be beneficial to mangroves, yet abrupt increases in sediment supply associated  
353 with rainfall and catchment land use change may be detrimental to mangrove forests.  
354 Extreme rainfall events in combination with changing land use practises may lead to  
355 increases in mangrove mortality due to smothering, yet may also increase mangrove  
356 resilience to sea level rise. We therefore highlight that synergistic impacts of anthropogenic  
357 activity and climate change may lead to unpredictable outcomes. Enhanced understanding of  
358 spatial variation in impacts and recovery using emerging remote sensing techniques (Lucas et  
359 al., 2017) will assist in managing coastlines exposed to increasing frequency of storms and  
360 climatic extremes.

361

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